# Application of Numerical Methods for New Estimate of Rheology Constants in the 2D Computer Model of the Mantle Wedge Thermal Convection as a Possible Physical Mechanism of Hydrocarbons Transport

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#### 9 Abstract

For both Newtonian and non-Newtonian mantle rheology laws, the numerical model of the 2D dissipationdriven mantle wedge thermal convection is constructed for the case of subduction of the Black sea micro-plate under the Crimea peninsula with the account taken of the phase transitions in the mantle. The horizontal extent of the positive 2D heat flux anomaly zone localized in the rear of the Crimea mountains is shown to correspond to the model subduction velocity ?10 mm per year for the water content of one weight

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*Index terms* — 2D thermal convection, Newtonian and non-newtonian rheology constants, phase transitions,
 hydrocarbons transport.

# <sup>19</sup> 1 Introduction

nteraction of the lithospheric plates in the Crimea-Caucasus region leads to the thrusting of the Black Sea microplate under the Crimea peninsula (under the Scythian plate) ??Nimetulayeva, 2004]. As a consequence, the
seismic focal plane is formed along which the Crimea ascends as the result of seismic jerks. The velocities of
vertical uplift of the Crimea mountains and sinking of the near-Crimean area of the Black Sea micro-plate equal
to ~4 mm per year and ~10 mm per year, respectively. Mountainous Crimea is a folded fault region being a part
of the Alps-Himalaya-Indonesia belt [Yudin, 2001].

In ??Ushakov et al., 1977] the subduction velocity of the Black Sea micro-plate under the Crimea peninsula is estimated of ~1 mm per year as the best fit to the observed sedimentary layer distribution. Other estimations are unknown to the knowledge of the authors. However the obtained estimate of ~1 mm per year appears to be an underestimate, being not correspondent to the vertical velocities of ~4 and ~10 mm per year of Mountainous Crimea and the Black Sea micro-plate.

- The abovementioned picture is reproduced here in Fig. 1 with the convective vortices drawn additionally. It is
- 42 worth assuming the two heat flux anomaly maxima observed in the south of the Crimea peninsula ??Smirnov,

According to [Gavrilov, 2014; ??erya, 2011; ??erya et al., 2006] two types of dissipation driven smallscale 31 thermal convection in the mantle wedge are possible, viz. 3D finger-like convective jets, raising to volcanic chain, 32 and 2D transversal Karig vortices, aligned perpendicularly to subduction. These two types of convection are 33 34 shown to be spatially separated due to the pressure and temperature dependence of mantle effective viscosity, 35 the Karig vortices, if any of them formed, being located behind the volcanic arc [Gavrilov, 2014]. Despite the 36 firmly established localization of the seismic focal plane there is just a single definite conclusion concerning the velocity of subduction of the Black Sea micro-plate ??Ushakov et al., 1977]. It is not completely clear if volcanism 37 played a substantial role in forming Mountainous Crimea, or the mountains are of a purely thrust-and-fold origin. 38 ??Nimetulayeva, 2004] indicates the contradictory statements on the Crimean volcanism to have been published, 39 however in Fig. 2.4 in ??Nimetulayeva, 2004], the volcanic eruption in the Mountainous Crimea is depicted. 40

1980; ??imetulayaeva, 2004, Fig.2.4] owe their origin to respectively 3D and 2D upward convective heat transfer 43 from the mantle wedge to the Earth's surface (see Fig. 1 of this paper). The latter 2D maximum located in 44 the rear of the Mountainous Crimea is much greater as compared to the former 3D maximum located in the 45 Mountainous Crimea. The 2D I © 2020 Global Journals lobal Journal of Researches in Engineering ( ) Volume 46 47 Xx X Issue I Ve rsion I heat flux anomaly maximum is associated with the 2D upward convective flow in the mantle wedge. Numerical modeling of 2D mantle wedge thermal convection occurring in the form of the Karig 48 vortices and presumably transporting heat to the Earth's surface in the rear of the Mountainous Crimea allows 49 judging about the mean velocity of subduction of the Black Sea micro-plate under the Crimea peninsula as well 50 as about the rheological mantle parameters. The horizontal extent of the 2D heat flux anomaly in the rear of the 51 Mountainous Crimea is shown to correspond to the mean subduction velocity >10 mm per year for the observed 52 subduction angle 15°. Numerical convection models accounting for the effects of phase transitions as well as the 53 pressure, temperature, and viscous stresses viscosity dependence fit in well with the heat flux observational data 54 in the case of non-Newtonian mantle rheology at the mean concentration of water in the mantle wedge of  $\sim 1$  wt. 55 %. 56

## 57 **2 II.**

# <sup>58</sup> 3 Algorithm and Computation Complexity

Here? is dynamic viscosity,? and indices denote partial derivatives with respect to coordinates x (horizontal), z (vertical) and time t,? is the Laplace  $z \ge V$ ??,  $z \ge V$ ???, (3)

## 67 4 RT pV E

Following ?? Trubitsyn & Trubitsyn, 2014] we assume the phase functions Where for "wet" olivine  $A=5.3\times10$ 72 15 s -1 , m=2.5, the grain size h =10 -1 -10 mm, b  $* = 5 \times 10$  -8 cm is the Burgers vector [Zharkov, 2003], E 73 \* =240 kJ . mol -1 is activation energy, V \* =5 $\times$ 10 3 mm 3. mol -1 is activation volume, ? =300 GPa is 74 75 the shear modulus normalizing factor, R is the gas constant. At the chosen constants and the grain size h=1.676 mm, non-dimensional viscosity also denoted To check as to how the estimate of the velocity of subduction of 77 the Black Sea micro-plate is sensitive to the accepted linear rheological law here we make extra computations for non-Newtonian rheology, in which case the viscosity formulae (5)-Where according to ?? Trubitsyn, 2012] for 78 "wet" olivine n=3, r=1.2, m=0,) (l ? as ? ? ? ? ? ? ? ? ? ? ? ? ? ? ) () () ( 1 2 1 1 1 1 w T z z th ,) (? ? ) ( 79 ) (0) () (0) (1111 T T g z T z ??? , (10)2 / 1 2) ? (? ik = E = 480 kJ . mol - 1, 80 V \* =11×10 3 mm 3. mol -1 , A=10 2 ? -1 ×(MPa) -n , C w >10 -3 for "wet" olivine is the weight water

81 V \* =11×10 3 mm 3. mol -1 , A=10 2 ? -1 ×(MPa) -n , C w >10 -3 for "wet" olivine is the weight water 82 concentration (in %%).

It should be noted the constants in (7) vary considerably in the papers referred to by ??Trubitsyn, 2012] and heretofore, we gave averaged values of constants. At C w =10 -3 on accounting for where the signs are changed as z-axis is pointing upwards, ) () (

## 86 5 T z l

87 ??????????)()(0)()(02)()(?/)(??2?,(**11**) x11111111111111 x T w T T Ra z z ch w Ra 88 ?????????)()(0)()()()(02)()()()()()()()(????2?????.(**12**)

Equations (1)-(2) are solved for the isothermal horizontal and insolated vertical boundaries regarded noslip impenetrable ones except for the "windows" for in-and outgoing subducting plate, where the plate velocity is specified. Vertical boundary distant from subduction zone is assumed penetrable at right angle, the latter boundary condition appears not too imposing in the case of very flat subduction. Q in (2) is non-zero in the continental and oceanic crust 40 and 7 km thick. Initial vertical boundaries temperature is calculated for the half-space cooling model for 10 9 yr and 10 8 yr for Scythian (continental) and Black Sea (oceanic) plates respectively.

## 96 6 III.

#### **7** Results and Discussion

Assuming the second (more remote from the trench) heat flux q maximum in Fig. 1 appears above the convective flow, ascending to C 2 point in Fig. 1, and the convection cell dimension is equal to the two adjacent q minima separation (i.e. the q minima are located above the descending convective flows) we can estimate the convection cell dimension as ~250 km. To preliminarily access the mean velocity of subduction of the Black Sea micro-plate the coordinate

103 x dependence of the growth rate ?? ( x ) of transversal convective rolls for the constant viscosity fluid model 104 can be allowed for. In such the model the averaged temperature and pressure viscosity dependence is accounted 105 for in an averaged manner, the factor describing the temperature-and pressure viscosity dependence being equal 106 to its mean value [Gavrilov, 2014].

Analytical formulae in [Gavrilov, 2014] It should be noted the growth rates ?? (x) are viscosity independent as convection is driven by viscous heat release (which is directly proportional to viscosity), while, on the other hand, the greater is the viscosity the more difficult is to arouse the convection. Fig. 2 clearly demonstrates the convective zone with ?? (x)>0 amounts to ?? 1 2 x x

111 250 km (i.e. the single convective cell of ?250 size is actually aroused) at V=40.5 mm per year, the latter value 112 being a preliminary estimate of the mean subduction velocity. The ?? (x) maximum is ?320 km distant from 113 the trench which is very close to the distance from the trench to the observed 2D heat flux anomaly (?400 km, 114 see Fig. 1).

To compute more accurate consistent model of small-scale convection in the mantle wedge between the 115 overriding Scythian plate and subducting Black Sea micro-plate it is necessary from the computational point 116 of view first to specify vanishing non-dimensional numbers Ra ?0, Di =0 in (1)-(2), i.e. to ignore convection 117 and viscous dissipation. This approach is applied as convection with Ra and Di (4) passes through very vigorous 118 119 stages, and the time steps in integrating (1)-(2) become too small thus making it difficult to model the thermal structure of the plates. Solving (1)-(2) by the finite element method in space on the grid  $104 \times 104$  and the 3-rd 120 order Runge-Kutta method in time one obtains for Ra ?0, Di =0 and V=45 mm a year non-dimensional quasi 121 122 steady-state? and T shown in Figs. 3, where the streamlines are depicted with step 0.25 and the isotherms with an interval of 0.05. Subducting plate was considered rigid, while the viscosity at the zone of plates friction 123 (at temperatures below 1200 K) was reduced by 2 orders of magnitude as compared to (5). The latter viscosity 124 reduction at the plates contact zone accounts for lubrication effected by deposits partially entrained by the 125 126 subducting plate. Such a lubrication prevents the overriding Scythian plate from gluing to the subducting one ??Gerya, 2011]. It is worth noting the isotherm T=0.15 in Fig. 3a,c approximately corresponding to the Earth's 127 128 surface is depressed at subduction zone by~7 km which is of the order of a typical trench depth. Fig. 3 shows 129 the results of computation for formulae (7) -(9) for non-Newtonian rheology case for the water content C w =10 130 -3 weight %% (Fig. 3a, b) and C w =  $3 \times 10$  -1 weight %% (Fig. 3c, d). The velocity V= 45 mm per year is chosen as resulting in the best convective zone size fitting in with the observed heat flux (positive and negative) 131 anomaly size at the point C 2 in Fig. 1, i.e. in the rear of the Mountainous Crimea. The Black Sea microplate 132 subducting with a given velocity V is considered rigid and is shown in Fig. 3b,d by the equidistant diagonal 133 streamlines. The induced mantle wedge flow above the subducting plate is seen to occur in the form of a single 134 vortex at C w = 10 - 3 weight %% (Fig. 3b These convective vortices are seen actually to correspond to a single 135 convection cell aroused at subduction velocity V=45 mm per year. The latter convection cell dimension is of the 136 order of ~300 km, i.e., is very close to the observed minima q separation under the C 2 point in Fig. 1. 137

138 Thus the for the non-Newtonian mantle wedge rheology case with the viscosity reduced by 3 orders of magnitude 139 as compared to (7)-9) the computation shows the convection in the mantle wedge to occur at C w  $=3\times10$  -1 weight %% in the form of two micro vortices at V=45 mm per year. Convection of this type can provide abnormal 140 2D heat flux q observed in the rear of the Mountainous Crimea and the upwelling of the mantle hydrocarbons 141 to the Earth's surface along the arrow "c" [Yudin, 2003]. Considerable velocity in convective vortices in Fig. 4 142 is due to the local viscous stresses increase resulting in the drop in viscosity in convective zone. In the case of 143 Newtonian rheology the convection is aroused at the subduction velocity of over  $10.2 \text{ mm} \times a - 1$ , which appears 144 unrealistic. 145

According to ??Zharkov, 2019, p.143], the water content in the mantle transition zone in the mantle wedge 146 may amount to  $\sim 3$  wt. %. To investigate the role of water infused into the mantle wedge from the subducting 147 slab the above computations were carried out for the mean water content of 1 wt. % and subduction velocity of 148 30, 20, and 10 mm per year. The results of the convection computation are shown in Figs. 5a and 5b for V=30 149 150 and 20 mm per year respectively, where the streamlines corresponding to subducting Black Sea micro-plate are 151 shown with the interval of 10, and the streamlines, corresponding to convective vortices with the interval of 10 6 152 . The mean non-dimensional velocity in the left micro-vortex are  $\sim 15.2 \times 10$  7,  $\sim 7.1 \times 10$  7 and  $\sim 0.05 \times 10$  7 for the velocity of subduction of V = 30, 20, and 10 mm per year respectively. Thus, the convection may be considered 153 to arise at the subduction velocity over  $\sim 10$  mm per year for the mean water content C w  $\sim 1$  wt.%. Since the 154 meant water content in the mantle wedge could hardly exceed ~1 wt.% even at the water content in the mantle 155 transition zone of 3 wt%, the obtained subduction velocity of ~10 mm per year may be regarded the minimum 156 estimate of that of subduction of the Black Sea micro-plate. 157

It is worth noting, that in the case of Newtonian rheology, the mantle wedge dissipation-driven convection in 158 the form of transversal rolls, as in Fig. 4, is characteristic of very small subduction angles, the convection of this 159 type being absent already at subduction angle ?=30° [Gavrilov & Abbott, 1999]. At the subduction angle under 160 consideration here,  $?=15^{\circ}$ , the convective transversal rolls do not appear at V<10 cm×yr -1 for the Newtonian 161 rheology case. Arrow (c) above the boundaries of the oppositely revolving convective vortices in Figs.4, 5 indicate 162 a possible direction of transport of non-organic mantle hydrocarbons to the Earth's surface. Computations for 163 Newtonian mantle rheology with the viscosity (5)-(6) shows the transversal rolls to be aroused at far greater 164 distance from the trench than the observed 2D heat flux anomaly. the model constructed here favors the non-165 Newtonian mantle wedge rheology as better fitting in with the observed heat flux anomaly localization. It should 166 be noted that numerous thermo-mechanical mantle models in the zones of subduction (see, e.g. ??Gerya et al., 167 2006; ??erya, 2011] and the vast number of references there) showed convection in the form of transversal rolls 168 never to occur as the models with extremely small subduction angle and sufficiently great subduction velocity 169 were not investigated. 170

171 IV.

#### 172 8 Conclusions

<sup>173</sup> The size of the cell of 2D mantle wedge dissipation-driven convection in the case of the realistic non-Newtonian rheology equals  $\sim 300$  km at the subduction velocity 10 mm×yr <sup>1 2</sup>



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Figure 1:

<sup>&</sup>lt;sup>1</sup>Application of Numerical Methods for New Estimate of Rheology Constants in the 2D Computer Model of the Mantle Wedge Thermal Convection as A Possible Physical Mechanism of Hydrocarbons Transport

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Figure 2:



Figure 3:



Figure 4:



Figure 5: T



Figure 6: 3 T



Figure 7: Fig. 1 :



Figure 8:

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